

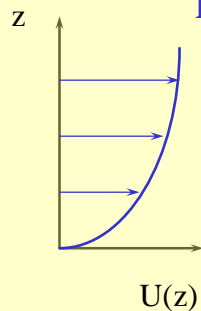
# Advanced Dynamical Meteorology

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CH 04

## Energetics of waves on stratified shear flows

**Boussinesq fluid**



**Wave energy equation (see Ex. 3.4)**

$$\frac{\partial E}{\partial t} = -\frac{\partial F}{\partial z} - \overline{\rho u w} U_z$$

**E = mean wave energy density**

**F = mean rate of working of the disturbance pressure force in the vertical =  $\overline{p w}$**

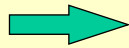
$$F = \overline{pw}$$

In the case of a non-moving medium ( $U = 0$ ),  $F$  is interpreted as a **mean flux of wave energy** and equals  $Ew_g$  (Ex. 2.11).

It is tempting to retain this interpretation of  $F$  in a moving fluid and to regard the term  $U_z$  in

$$\frac{\partial E}{\partial t} = -\frac{\partial F}{\partial z} - \rho \overline{uw} U_z$$

as a '**source**' of mean wave energy associated with the interaction of the wave with the basic shear,  $U_z$ .



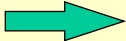
This interpretation can be misleading!

Question: What is meant by '**energy flux**' in a moving medium?

To see why  $-\rho \overline{uw} U_z$  in  $\frac{\partial E}{\partial t} = -\frac{\partial F}{\partial z} - \rho \overline{uw} U_z$  ought not

to be interpreted as an energy source, consider the equation for the mean flow to second-order in wave amplitude.

Full horizontal momentum equation in flux form



$$\rho u_t^* + (\rho u^{*2})_x + (\rho u^* w)_z = -p_x$$

$u^*$  is the **total** horizontal wind speed

Put  $u^* = \bar{u}(z, t) + u(x, z, t)$

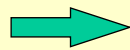
$u$  = **deviation** from the mean wind  $\bar{u}$

**mean** refers to an average over a wavelength:  $\overline{(\ )} = \frac{1}{\lambda} \int_0^\lambda (\ ) dx$

or for an non-periodic disturbance which vanishes as  $x \rightarrow \pm\infty$ :

$$\overline{(\ )} = \int_{-\infty}^{\infty} (\ ) dx$$

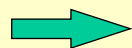
Substitute for  $u^*$  in  $\rho u_t^* + (\rho u^{*2})_x + (\rho u^* w)_z = -p_x$  and average



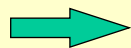
**the mean flow momentum equation**

$$\rho \bar{u}_t + (\rho \bar{u} w)_z = 0$$

Assume that the wave amplitude is sufficiently small



$$\bar{u} = U(z) \quad \text{and} \quad \bar{w} = 0$$



**the interaction between the perturbation and the basic flow can be ignored.**

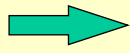
However, the waves have a second-order effect in amplitude on the mean flow governed by

$$\rho \bar{u}_t + (\rho \bar{u} w)_z = 0 \quad \text{(see ADM)}$$

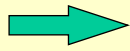
$\times \bar{u} \Rightarrow$  **the mean flow kinetic energy equation**

$$\frac{\partial}{\partial t} \left( \frac{1}{2} \rho \bar{u}^2 \right) = -\bar{u} (\rho \bar{u} w)_z$$

$$\frac{\partial}{\partial t} \left( \frac{1}{2} \rho \bar{u}^2 \right) = -\bar{u} (\rho \bar{u} w)_z$$



local second-order changes in the mean flow are associated with nonzero values of  $\bar{u} (\rho \bar{u} w)_z$



this term should appear as the "source term" in

$$\frac{\partial E}{\partial t} = -\frac{\partial F}{\partial z} - \rho \bar{u} w U_z$$

Rewrite as

$$\frac{\partial E}{\partial t} = -\frac{\partial}{\partial z} (\bar{p} w + U \rho \bar{u} w) + U (\rho \bar{u} w)_z$$

Interpret

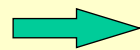
$F = \bar{p} w + U \rho \bar{u} w$  as the total or net energy flux.

Not Galilean invariant

### Perturbation analysis

Put  $u = \varepsilon u_1 + \varepsilon^2 u_2 + \dots$  and  $\bar{u} = U(z) + \varepsilon^2 \bar{u}_2(z, t) + \dots$

where  $\varepsilon \ll 1$  and the  $u_i$  etc. are  $O(1)$ , in  $\rho \bar{u}_t + (\rho \bar{u} w)_z = 0$



$$\rho \bar{u}_{2t} + (\rho \bar{u}_1 w_1)_z = 0$$

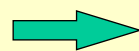
Assumes that the wave-induced vertical motion is zero at  $O(\varepsilon^2)$

Wave energy equation

Mean flow kinetic energy

$$\frac{\partial E}{\partial t} = -\frac{\partial}{\partial z} (\bar{p} w + U \rho \bar{u} w) + U (\rho \bar{u} w)_z \quad \frac{\partial}{\partial t} \left( \frac{1}{2} \rho \bar{u}^2 \right) = -\bar{u} (\rho \bar{u} w)_z$$

At  $O(\varepsilon^2)$



$$\frac{\partial}{\partial t} (E + \rho U \bar{u}_2) = -\frac{\partial}{\partial z} (\bar{p}_1 w_1 + U \rho \bar{u}_1 w_1)$$

$$\frac{\partial}{\partial t}(\underline{E} + \rho U \bar{u}_2) = -\frac{\partial}{\partial z}(\overline{p_1 w_1} + U \rho \overline{u_1 w_1})$$

**wave energy plus  
mean flow energy**

$$\rho U \bar{u}_2 = \frac{1}{2} \rho (U + \varepsilon^2 \bar{u}_2 + \dots)^2 - \frac{1}{2} \rho U^2$$

**to  $O(\varepsilon^2)$**

**the divergence of the vertical advective  
flux of **total** kinetic energy**

$$\frac{1}{2} \rho (U + \varepsilon u_1 + \dots)^2 \text{ at } O(\varepsilon^2)$$

**= the divergence of the first nonzero  
term in the expression**

$$\frac{1}{2} \rho (U + \varepsilon u_1 + \dots)^2 \varepsilon w_1$$

### Exercise

Show that the **perturbation and mean flow** equations:

$$\frac{\partial E}{\partial t} = -\frac{\partial F}{\partial z} - \rho \overline{u w} U_z \qquad \frac{\partial}{\partial t} \left( \frac{1}{2} \rho \bar{u}^2 \right) = -\bar{u} (\rho \overline{u w})_z$$

form an **energetically closed system** in the sense that, for a free wave with

$$F(0) = 0 \quad \text{and} \quad F \rightarrow 0 \text{ as } z \rightarrow \infty$$

**recall that**  
 $F = \overline{p w}$

$$\int_0^\infty \left( E + \frac{1}{2} \rho \bar{u}^2 \right) dz \quad \text{is a constant}$$

## The nonacceleration theorem

The perturbation equations for waves in a Boussinesq fluid are:

$$u_t + Uu_x + wU_z = -P_x$$

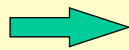
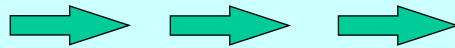
$$w_t + Uw_x = -P_x + \sigma$$

$$\sigma_t + U\sigma_x + N^2w = 0$$

$$u_x + w_z = 0$$

Look for steady travelling wave solutions of the form

$$(u, w, P, \sigma) = \text{Re}\left[(\hat{u}(z), \dots)e^{ik(x-ct)}\right]$$



$$ik(U - c)\hat{u} + U_z\hat{w} = -ik\hat{P}$$

$$ik(U - c)\hat{w} = -\hat{P}_z + \hat{\sigma}$$

$$ik(U - c)\hat{\sigma} + N^2\hat{w} = 0$$

$$ik\hat{u} + \hat{w}_z = 0$$

Put  $V = U - c$

$$\hat{\sigma} = \frac{iN^2}{kV}\hat{w}$$

$$\hat{P} = -\frac{i}{k}(V\hat{w}_z - V_z\hat{w})$$

$$\hat{P}_z = \left[-ikV + \frac{iN^2}{kV}\right]\hat{w}$$

$$\hat{w}_{zz} + \left[\frac{N^2}{V^2} - \frac{V_{zz}}{V} - k^2\right]\hat{w} = 0$$

$$\hat{w}_{zz} + \left[ \frac{N^2}{V^2} - \frac{V_{zz}}{V} - k^2 \right] \hat{w} = 0$$

This is the Boussinesq form of Scorer's equation

We can write the components of

$$(u, w, P, \sigma) = \text{Re}[(\hat{u}(z), \dots)e^{ik(x-ct)}]$$


as  $u = \frac{1}{2} \hat{u}(z)e^{ik(x-ct)} + (\text{---})^*$  etc.

**complex conjugate**

Then, for any two dependent variables a and b

$$\overline{ab} = \frac{1}{4}(\hat{a}\hat{b}^* + \hat{a}^*\hat{b})$$


**Multiply**  $ik(U - c)\hat{u} + U_z \hat{w} = -ik\hat{P}$  **by**  $\rho \hat{w}^* / ik$

**and use**  $ik\hat{u} + \hat{w}_z = 0$  

$$(U - c)\rho \hat{u} \hat{w}^* - \frac{i}{k} \rho U_z |\hat{w}|^2 = -\hat{P} \hat{w}^*$$

**Add this equation to its complex conjugate and use**

$$\overline{ab} = \frac{1}{4}(\hat{a}\hat{b}^* + \hat{a}^*\hat{b})$$

  $(U - c)\rho \overline{u\hat{w}} = -\overline{P\hat{w}}$

Also, for wave perturbations such as

$$(u, w, P, \sigma) = \text{Re}[(\hat{u}(z), \dots)e^{ik(x-ct)}]$$

$$F = \overline{pw} + U\rho\overline{uw} = \frac{cF}{c-U} = c\rho\overline{uw}$$

For a steady (in amplitude) sinusoidal wave  $\partial E/\partial t = 0$  and from

$$\frac{\partial E}{\partial t} = -\frac{\partial F}{\partial z} - \rho\overline{uw}U_z \quad \text{and} \quad (U-c)\rho\overline{uw} = -\overline{pw}$$



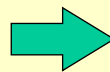
$$\frac{\partial F}{\partial z} = -\rho\overline{uw}U_z$$



$$(U-c)\rho\overline{uw} = -F$$

$$(U-c)\frac{d}{dz}(\rho\overline{uw}) = 0$$

$$(U-c)\frac{d}{dz}(\rho\overline{uw}) = 0$$



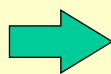
$$U = c$$

or

$$\frac{d}{dz}(\rho\overline{uw}) = 0$$

$$\frac{d}{dz}(\rho\overline{uw}) = 0 \quad \text{and} \quad F = c\rho\overline{uw} \quad \Rightarrow \quad \frac{dF}{dz} = 0$$

$$\rho\overline{u}_t = -(\rho\overline{uw})_z$$



the waves do not force any second-order acceleration of the mean flow



## The nonacceleration theorem

For a steady (in amplitude) sinusoidal wave

$$\frac{d}{dz}(\rho \overline{uw}) = 0 \quad \text{and} \quad \rho \overline{u_t} = -(\rho \overline{uw})_z \quad \Rightarrow \quad \overline{u_t}$$

This result is now known as the **nonacceleration theorem**.

It was first obtained in a slightly less general form by **Eliassen and Palm (1960)** and has been shown to be a quite general result by **Andrews and McIntyre (1978)**.

The quantity  $F = c\rho \overline{uw}$  is the **total vertical flux of wave energy**.

## Flow over sinusoidal orography

For waves which radiate vertically ( $0 < |k| < 1$ ), there exists a downward flux of mean horizontal momentum ( $\rho \overline{uw} < 0$ ).

$\rho \overline{uw} < 0$  is independent of height and equal to the drag per unit wavelength exerted by the boundary on the airstream (see **Ex. 3.5**).

$$\rho \overline{uw} = -\frac{1}{2} \rho U^2 m k h_m^2$$

where  $\text{sgn}(mk) > 0$  for upward propagation.

Evidently, the momentum flux originates at infinity, where, presumably, the drag exerted by the boundary on the air stream acts. **This is at first sight puzzling !**

When a general airstream  $U(z)$  flows over mountain ridge and produces **upward radiating waves**, a **forward wave drag** is exerted on the mountain.



The mountain exerts a drag on the airstream

**Question:** How is the stress on the airstream distributed? i.e. at what level(s) does the drag act on the airstream?

**For steady waves**, the nonacceleration theorem rules out the possibility of interaction except at a **critical level** where  $U = c$ , a level where the intrinsic frequency of the waves vanishes (see Eq. 3.4).

In the case of **stationary mountain waves**,  $c = 0$ .

- Linear theory suggests that **at a critical level, the wave is almost completely absorbed**, leading to a **deceleration** of the mean flow at that level.
- However, nonlinear and viscous effects may be important near the critical level.
- I shall not address the **critical-layer problem** in this course - for further details, see the important paper by **Booker and Bretherton (1967)**.
- For a propagating wave packet with a spectrum of horizontal phase speeds, there may be a **range of critical levels**.
- Then absorption by the mean flow takes place in a finite layer.
- For stationary mountain waves,  $c = 0$  for all Fourier components.

The analysis of **unsteady wave development** and of **wave propagation in arbitrary shear flows** is mathematically difficult.

For **unsteady wave development** Laplace transforms methods for initial-value problems usually lead to uninvertible transforms.

For **wave propagation in arbitrary shear flows**, typical eigenvalue problems are analytically intractable, or at best, very complicated.

Some analytical progress and further physical insight may be obtained by studying slowly-modulated wave trains or wave packets, in which the waves are locally plane, with wavelength and amplitude varying significantly only over a space scale of many wavelengths

### Slowly varying wave trains or wave packets

Theory of **Acheson** (1976, pp 452 – 455):

Assume the waves have constant frequency  $\omega$  and horizontal wavenumber  $k$ , but their **amplitude varies with height and time on scales very long compared with one wavelength and one period, respectively.**

Define '**slow**' variables:  $Z = \alpha z$  and  $T = \alpha t$ , where  $\alpha \ll 1$  is a dimensionless measure of how slowly the wave train is modulated.

**Slowly varying** means that at any given height/time the wave amplitude varies by a factor  $O(1)$  over a height/time scale of  $O(\alpha^{-1})$  wavelengths/periods.

Assume a **Boussinesq fluid** with  $N$  constant, but with a basic shear  $(U(Z), 0, 0)$ .

Consider linear wave perturbations with streamfunction

$$\psi = \psi_1 + \alpha \psi_2 + \dots$$

where  $\psi_n = \text{Re} \left[ \hat{\psi}_n(Z, T) e^{i(kx + \theta(z) - \omega t)} \right]$

Similar expansions are taken for other flow quantities.

A local vertical wavenumber defined in terms of the phase function  $\theta(z)$  is

$$m(Z) = \frac{d\theta}{dz}$$

### The multiple-scaling technique

Replace time and height derivatives by

$$\frac{\partial}{\partial t} \rightarrow \frac{\partial}{\partial t} + \alpha \frac{\partial}{\partial T} \quad \text{and} \quad \frac{\partial}{\partial z} \rightarrow \frac{\partial}{\partial z} + \alpha \frac{\partial}{\partial Z} \quad \rightarrow$$

$$u_t + U u_x + P_x = -(\alpha u_T + w \underline{U_Z})$$

$$w_t + U w_x + P_z - \sigma = -(\alpha w_T + w P_Z)$$

$$\sigma_t + U \sigma_x + N^2 w = -\alpha \sigma_T$$

$$u_x + w_z = -\alpha w_Z$$

Term describing the effect of vertical shear appears only at  $O(\alpha)$ .

## Zero-order solution

Substitution of the expansions for  $u, w, P$ , gives to  $O(\alpha^0)$  a locally-plane wave solution, identical with that which would be obtained if  $U$  were a constant.

For this solution

$$\omega^{*2} = \frac{N^2 k^2}{k^2 + m^2}$$

$\omega^*(Z) = \omega - kU(Z)$  is the intrinsic frequency.

$\omega^*$  is a function of  $Z$  through  $U(Z)$  and  $m(Z)$ .

## First-order solution

At  $O(\alpha)$  in the expansion, the equations for subscript '2' quantities become:

$$i(\omega^* \hat{u}_2 - k\hat{P}_2) = \hat{u}_{1T} + U_Z \hat{w}_1$$

$$i(\omega^* \hat{w}_2 - m\hat{P}_2) + \hat{\sigma}_2 = \hat{w}_{1T} + \hat{P}_{1Z}$$

$$i\omega^* \hat{\sigma}_2 - N^2 \hat{w}_2 = \hat{\sigma}_{1T}$$

$$i(k\hat{u}_2 + m\hat{w}_2) = -\hat{w}_{1Z}$$

Eliminating  $\hat{u}_2, \dots$  etc. we obtain an expression of the form

$$(\dots)\hat{w}_2 = \text{expression involving T and Z derivatives of subscript '1' quantities}$$

where the coefficient of  $\hat{w}_2$  is zero.

$(\dots)\hat{w}_2 =$  [expression involving T and Z derivatives of subscript '1' quantities]



This expression, when set equal to zero, gives a solvability condition for the  $O(\alpha)$  problem.

This 4 \* 4 set of linear equations has the general form  $Ax = b$ , where A is a 4 \* 4 matrix with  $\det A = 0$  and x and b are column vectors.

It has a non-unique solution, but only if b is orthogonal to the solution y of the adjoint homogenous problem  $A'y = 0$ .

## Wave action

After some algebraic manipulation (see Appendix to ADM), the solvability condition for  $(\dots)\hat{w}_2 = \dots$  may be written in the form

$$\frac{\partial A}{\partial T} + \frac{\partial}{\partial Z}(w_g A) = 0$$

where

$A = E / \omega^*$  called the wave action

$E = \frac{1}{2} \rho_o (1 + m^2 / k^2) |\hat{w}_1|^2$

$w_g = \frac{\partial \omega}{\partial m}$  is the local vertical component of the group velocity

$\frac{\partial A}{\partial T} + \frac{\partial}{\partial Z}(w_g A) = 0$  expresses the conservation of wave action

### Exercise

Assuming that the mean second-order perturbation to the basic flow forced by the foregoing slowly-varying wave varies only with  $Z$  and  $T$ , the mean horizontal momentum equation may be written

$$\alpha(\partial \bar{u}_2 / \partial t + U_Z \bar{w}) = -(\overline{u_1 w_1})_Z$$

where, from continuity  $\bar{w}_Z = 0$

Show that 
$$\rho \frac{\partial \bar{u}_2}{\partial T} = -k \frac{\partial}{\partial Z} (w_g A)$$

Moreover, show that if  $\bar{u}_2 = 0$  in the absence of waves (i.e., when  $A = 0$ ), then

$$\rho \bar{u}_2 = \frac{E}{c - U}$$

$$\frac{\partial A}{\partial T} + \frac{\partial}{\partial Z} (w_g A) = 0 \quad \rightarrow$$

when the amplitude of the wave is independent of time,  $A w_g$  is independent of height.

$$\rightarrow \quad \overline{p_1 w_1} / (c - U) \quad \text{is independent of height.}$$


$$\frac{d}{dz} (\rho \overline{u w}) = 0 \quad \rightarrow$$


when the wave amplitude is steady, this is true even when no restriction is placed on how fast  $U$  varies over a vertical wavelength.

Consider now a wave train set up by the horizontal translation with speed  $c$  of a corrugated boundary at  $z = 0$ .

The corrugations are assumed to increase gradually in amplitude (from zero at time  $t$ ) on the slow time scale  $T$  as governed by the formula

$$\zeta(x,t;T) = A(T) \cos k(x - ct)$$

  $\bar{w} = 0$  at  $z = 0$

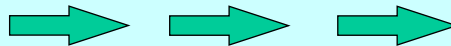
Thus  $\bar{w}_z = 0$    $\bar{w} \equiv 0$

As the slowly-modulated wave train propagates upwards past any given level, the local wave energy density  $E$  will slowly increase to a maximum and (in the absence of dissipation) then diminish again to zero as the wave train passes.

According to 
$$\rho \bar{u}_2 = \frac{E}{c - U}$$

the local modification  $\bar{u}_2$  to the mean flow varies similarly, the mean flow being accelerated (decelerated) if  $c > U$  ( $c < U$ ).

If the forcing  $\zeta(x,t;T) = A(T) \cos k(x - ct)$  slowly attains a constant amplitude  $A_0$  on the time scale  $T$  and persists at that amplitude thereafter, the wave train will consist of a precursor (which contains  $O(\alpha^{-1})$  wavelengths and whose amplitude increases with depth from effectively zero to that amplitude  $A_0$  which the source ultimately attains) and a lower part of constant amplitude  $A_0$  extending all the way down to the source.





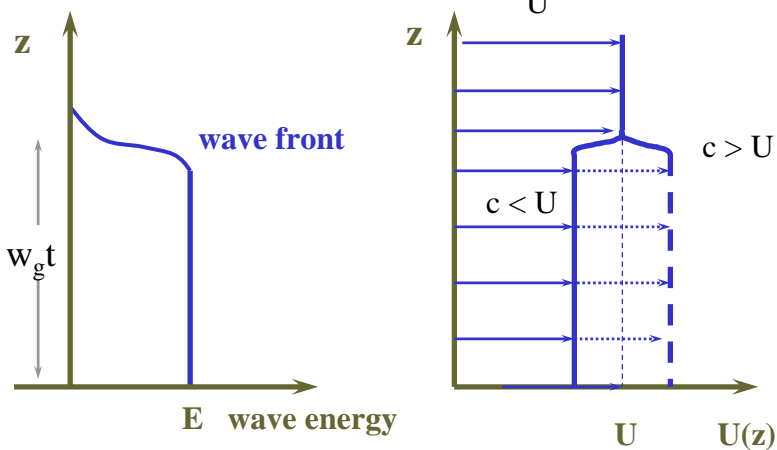
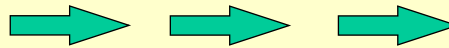
When  $U$  is a constant, so are  $m$  and  $w_g$  and

$$\frac{\partial A}{\partial T} + \frac{\partial}{\partial Z}(w_g A) = 0$$

reduces to the statement that amplitude modulations propagate upwards at the group velocity.

In particular, what we call for convenience the '**front**' of the wave train moves upwards at this speed.

the tolerably well-defined highest point at which the amplitude is  $A_0$



Wave energy density  $E$  as a function of height for a wave source switched on at  $z = 0$  at  $t = 0$  (**left**) and corresponding mean flow changes (**right**).

We may now understand the result that for steady flow over sinusoidal topography there is a downward flux of mean horizontal momentum **from infinity**.

If we imagine such a flow to be established by the gradual evolution of the topography as described by

$$\zeta(x,t;T) = A(T) \cos k(x - ct)$$

with  $c = 0$ , it is clear that the source of the downward momentum flux in the steady wave regime (i.e. at heights below  $z = w_g t$ ) is the deceleration of the mean flow in the region constituting the front of the wave train. **At no finite time is there a momentum flux at infinity.**

**The End**