

# 3DCATS – Synthetic data: impact of 3D clouds on retrieved NO<sub>2</sub> VCD

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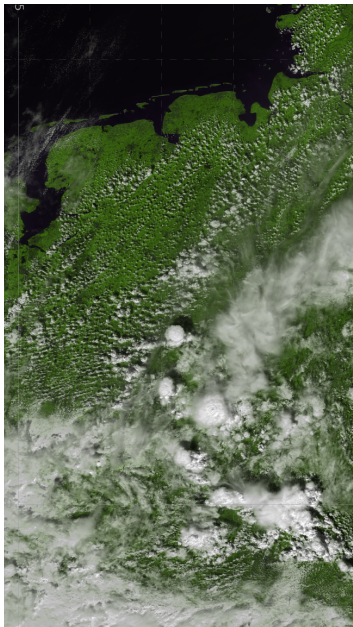
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# Introduction

- Operational retrievals of tropospheric trace gases from space-borne instruments based on 1D radiative transfer neglect
  1. cloud scattering into clear regions
  2. cloud shadows
- Monte Carlo radiative transfer (MYSTIC-ALIS)
  - ⇒ simulation of spectra for realistic 3D model atmospheres
- Application of  $\text{NO}_2$  retrieval algorithm on simulated data:
  - ⇒ estimation of retrieval error due to 3D cloud scattering



- **Aim:** Use synthetic data to validate and improve NO<sub>2</sub> retrieval algorithms
- **Radiative transfer model MYSTIC**
  - Horizontal photon transport is essential to investigate impact of cloud scattering on trace gas retrievals ⇒ Monte Carlo RT approach
- **Box cloud scenario**
  - Simulated spectra and layer-AMFs
  - investigate **sensitivity of NO<sub>2</sub> retrieval error** on various parameters
- **LES cloud scenario**
  - Cloud scene from ICON-LES model over Europe (698×763 km<sup>2</sup>)
  - All types of realistic clouds included
  - Representative sun-satellite geometries and surface albedos
  - Generate **synthetic dataset** for geostationary orbit and low Earth orbit for VIS and O<sub>2</sub>A-band
  - **Quantification of NO<sub>2</sub>-retrieval error**

# 3D radiative transfer in high spectral resolution

$\text{NO}_2$  retrieval (DOAS) – fit differential optical thickness

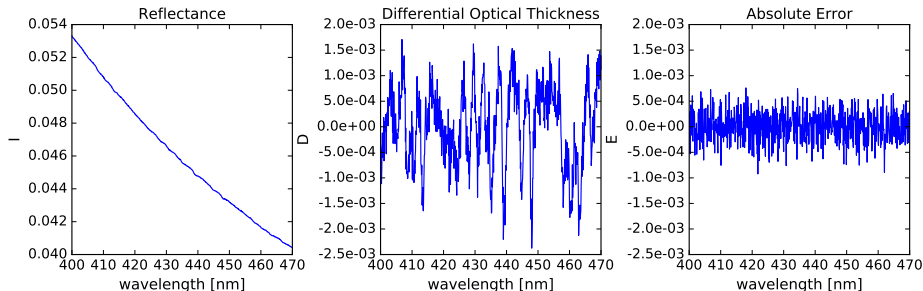
$$D(\lambda) = \ln(I_{TOA}(\lambda)) - P_3(\lambda)$$

$I_{TOA}$ : reflectance, spectral range:  $\lambda \approx 400\text{-}500$  nm

Radiative transfer requirements:

⇒ **high spectral resolution** (resolve characteristic absorption features)

⇒ **high accuracy** (absorption signal weak compared to Rayleigh continuum)



Standard Monte Carlo method: **computational time extremely high**

(**about 33h** for  $10^7$  photons/wavelength and 0.1 nm spectral resolution!)

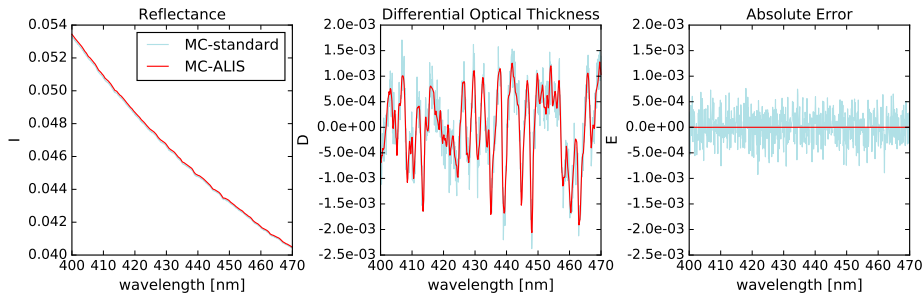
# Absorption Lines Importance Sampling

Trace photons at only one wavelength and calculate full line-by-line spectra

Spectral absorption and scattering included by photon weights

Statistical error causes bias (decreasing with  $\sqrt{N}$ ) over full spectral range, not for each wavelength

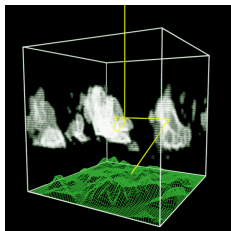
Computational time: **1.5 minutes** (comparable to DISORT)



C. Emde, R. Buras, and B. Mayer. *ALIS: An efficient method to compute high spectral resolution polarized solar radiances using the Monte Carlo approach*. JQSRT, 2011

# Radiative transfer model MYSTIC

*Monte carlo code for the phYSically correct Tracing of photons In Cloudy atmospheres (Mayer 2009)*

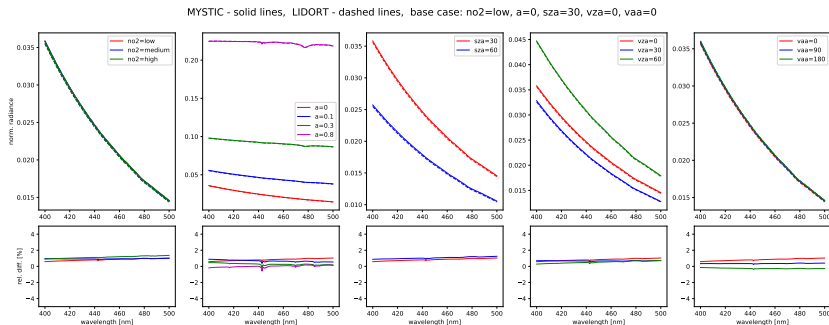


- *Special features:*

- Polarized radiative transfer (*Emde et al., 2010*)
  - **VROOM: variance reduction methods** (*Buras and Mayer, 2011*)
    - ⇒ radiance calculations for strongly peaked scattering phase functions
  - **ALIS method** (*Emde et al., 2011*)
    - ⇒ very efficient high spectral resolution calculations
  - complex topography (*Mayer et al., 2010*)
  - spherical geometry (*Emde and Mayer, 2007*)
  - **layer/box-airmass factors** in 3D domain (*Schwärzel, Emde et al. 2020*)
- Integrated in libRadtran package [www.libradtran.org](http://www.libradtran.org)  
(*Mayer and Kylling, 2005, Emde et al. 2016*)

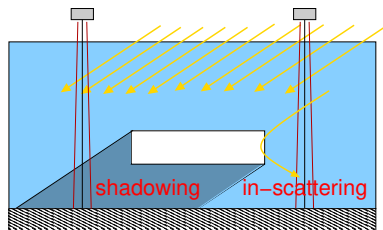
# Tests in one-dimensional geometry

- Simulation of spectra in VIS (400–500 nm, 0.2 nm resolution) and O<sub>2</sub>A-band (755–775 nm, 0.005 nm resolution)
- **Model intercomparison:** very good agreement between LIDORT and MYSTIC



- **Application of NO<sub>2</sub> retrieval algorithm** on synthetic spectra  
Clearsky  $\Rightarrow$  quantify retrieval error due to model differences  
1D cloud layer  $\Rightarrow$  test cloud correction schemes (O<sub>2</sub>-O<sub>2</sub> and FRESKO)

# Clearsky pixels in vicinity of clouds



Sketch of box cloud setup.

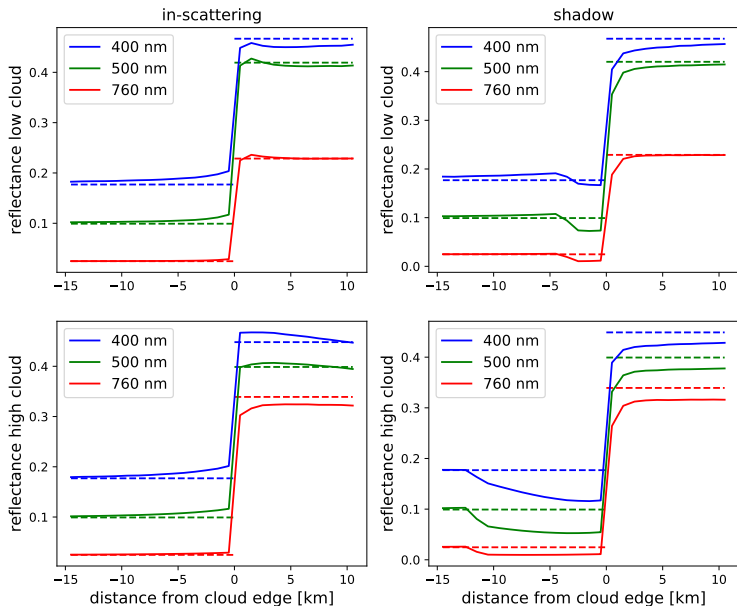
## General settings

- nadir observation geometry
- $1 \times 1 \text{ km}^2$  square field-of-view
- $\text{NO}_2$  profiles: Pacific polluted, European polluted
- surface albedo: 0.05
- no aerosol

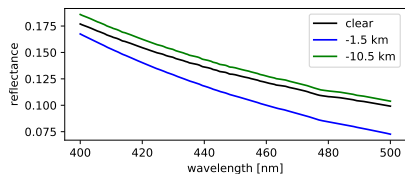
	liquid water cloud	ice water cloud
cloud optical thickness	1, 2, 5, <b>10</b> , 20	1, 2, <b>5</b> , 10, 20
cloud bottom height [km]	<b>2</b> , 5, 10	5, <b>9</b> , 12
effective radius [ $\mu\text{m}$ ]	10	30
optical properties	Mie	Baum (V3.6)
cloud geometrical thickness [km]	0.2, <b>1</b> , 2, 4, 8	
surface albedo	0.02, <b>0.05</b> , 0.1, 0.15, 0.2, 0.3	
solar zenith angle [ $^\circ$ ]	20, 30, 40, <b>50</b> , 60, 70, 80	



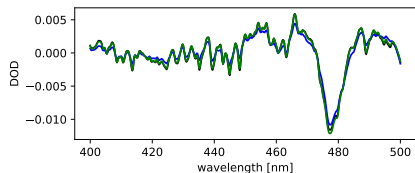
# Reflectance as function of distance from cloud edge



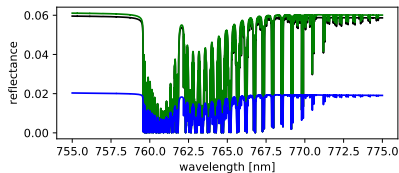
# Simulated spectra in clear region



- VIS: 400–500 nm,  $\Delta\lambda=0.2$  nm

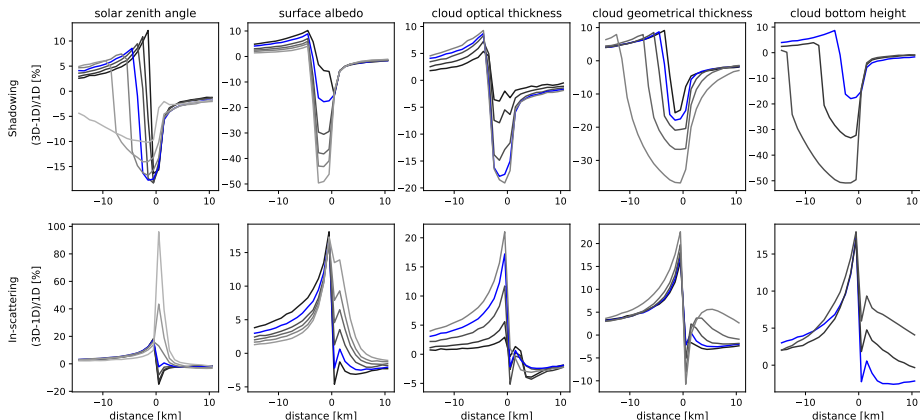


- $DOD(\lambda) = \ln(I(\lambda)) - P_3(\lambda)$



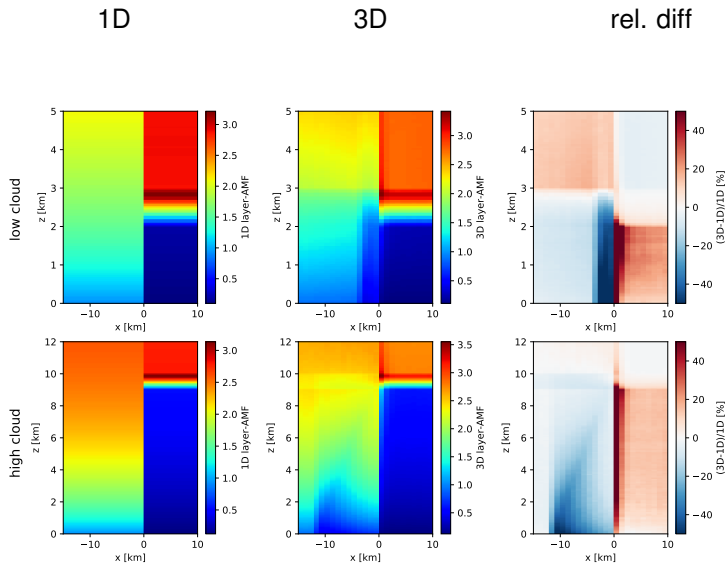
- O<sub>2</sub>A band: 755–775 nm,  $\Delta\lambda=0.005$  nm

# Rel. diff. between 3D and 1D reflectance (460nm)



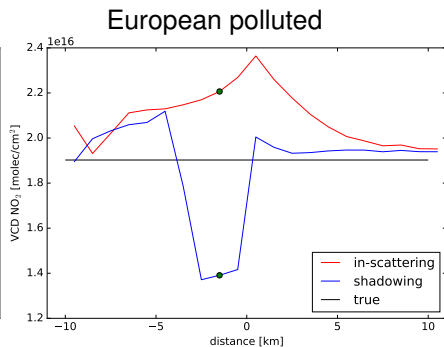
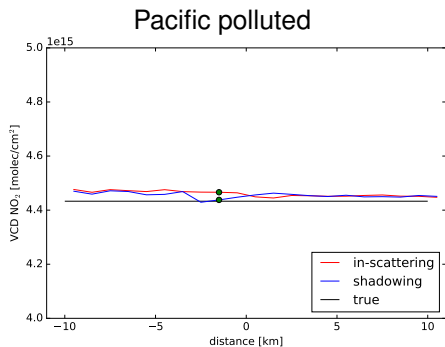
	liquid water cloud	ice water cloud
cloud optical thickness	1, 2, 5, <b>10</b> , 20	1, 2, <b>5</b> , 10, 20
cloud bottom height [km]	<b>2</b> , 5, 10	5, <b>9</b> , 12
cloud geometrical thickness [km]	0.2, <b>1</b> , 2, 4, 8	
surface albedo	0.02, <b>0.05</b> , 0.1, 0.15, 0.2, 0.3	
solar zenith angle [°]	20, 30, 40, <b>50</b> , 60, 70, 80	

# Layer-AMFs for base cases, cloud shadow



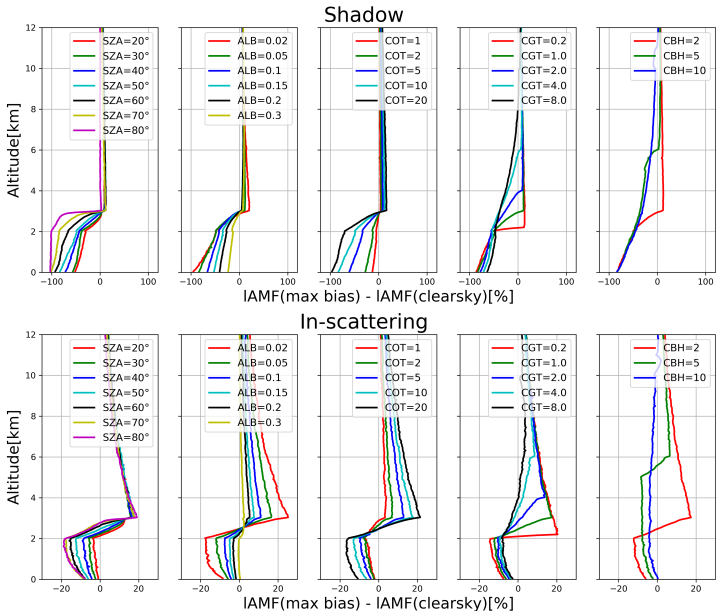
- **Standard approach:**
  - 1. DOAS fit is performed using the QDOAS software (Danckaert et al., 2017)  
⇒ NO<sub>2</sub> slant column densities (SCD)
  - 2. Convert SCD to vertical column densities (VCD) using air mass factors (AMF)
- **AMF calculation:** integral of the relative vertical NO<sub>2</sub> distribution, weighted by layer-AMF computed with a radiative transfer model (VLIDORT).
- **Cloud correction** uses cloud fraction, cloud top pressure and cloud top albedo from two cloud retrieval algorithms: O<sub>2</sub>-O<sub>2</sub> (Acarreta et al., 2004) and FRESCO (Koelemeijer et al., 2001)

# NO<sub>2</sub> retrieval results

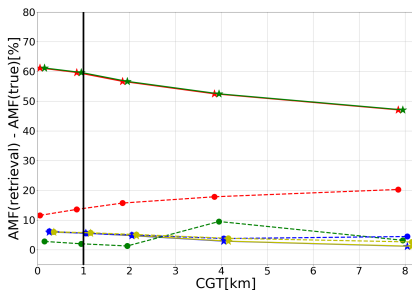
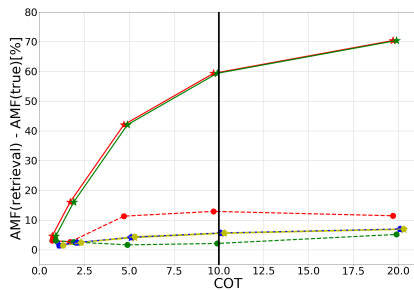
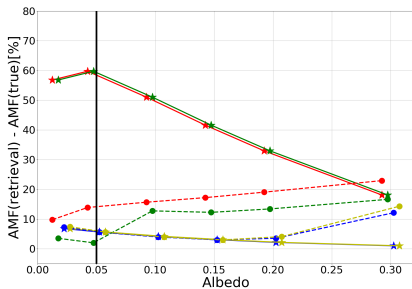
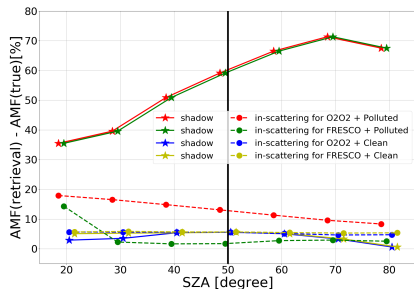


NO<sub>2</sub> retrieval results depending on distance from cloud edge.

# Layer-air-mass factor retrieval error

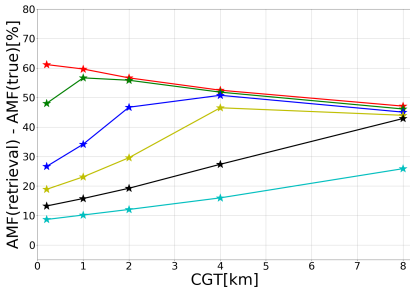
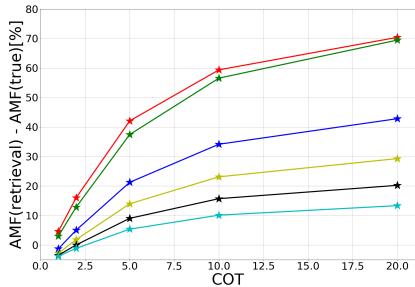
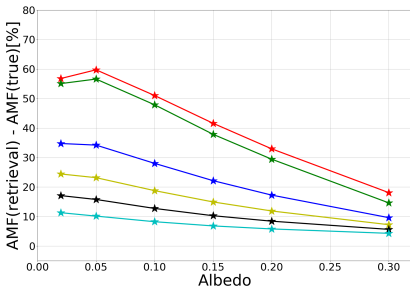
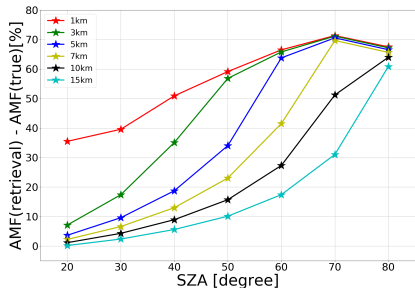


# Impact of NO<sub>2</sub> profile





# Impact of spatial resolution



# Summary - sensitivity of retrieval error

- retrieval error in cloud shadow is large
  - $\approx 35\%$  for  $\text{SZA}=20^\circ$  to more than  $70\%$  for  $\text{SZA}=80^\circ$
  - affected area extends to  $1\text{km}$  from the cloud edge for  $\text{SZA}=20^\circ$  to more than  $15\text{km}$  for  $\text{SZA}=80^\circ$
  - $60\%$  for low surface albedo ( $0.02/0.05$ ), decreases to  $10\%$  for  $\text{albedo}=0.3$
  - $5\%$  for  $\text{COT}=1$  to  $70\%$  for  $\text{COT}=20$
- retrieval error in in-scattering region relatively small ( $<20\%$ ), affected area is within  $5\text{ km}$  from cloud edge

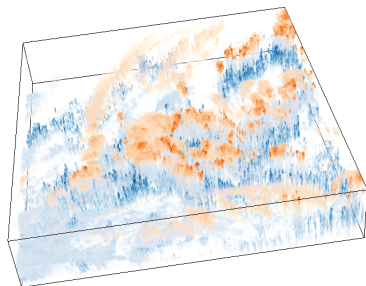
# LES cloud scenario

## ICON model

ICOsahedral Nonhydrostatic atmosphere model;

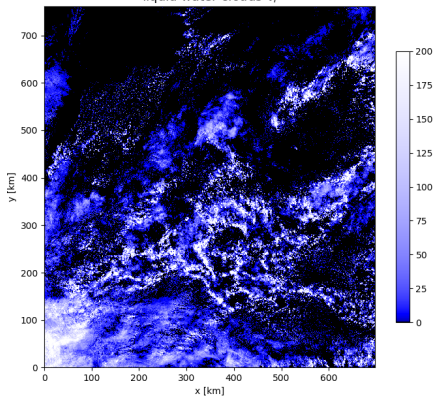
Dipankar et al. 2015, Zängl et al. 2015

- Spatial resolution approx. 1 km for region including Germany, Netherlands and parts of other surrounding countries
- Model validated against ground-based and satellite based observational data (Heinze et al. 2017)
- Simulations include all cloud types that are typical for Europe (e.g. shallow cumulus, cirrus, stratus, and convective clouds)

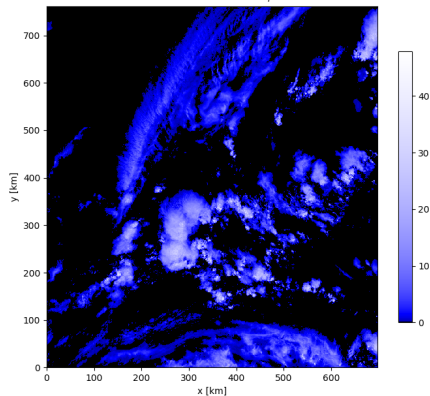


# Vertically integrated cloud optical thickness

liquid water clouds  $\tau_l$



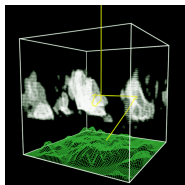
ice water clouds  $\tau_i$



# Reflectance simulation with LES clouds

## MYSTIC – Monte Carlo radiative transfer model

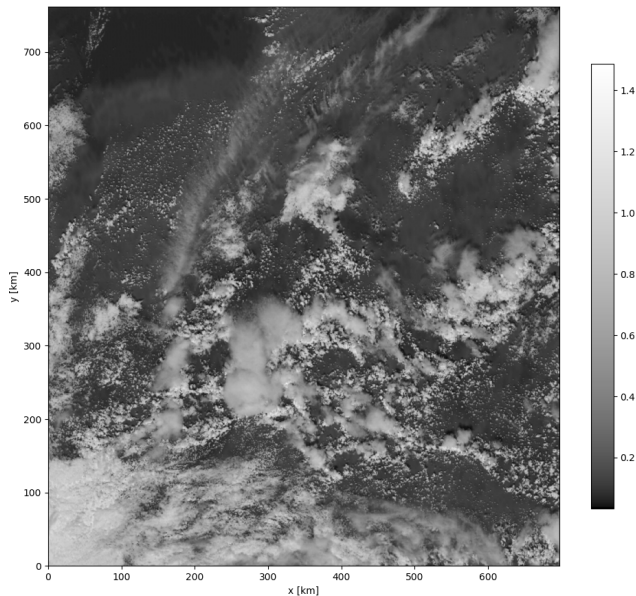
Mayer 2009, Emde et al. 2011, Emde et al. 2016



- Central wavelength 554 nm, Bandwidth 19.26 nm (Sentinel3 SLSTR B1)
- Nadir view, spatial resolution 1.2 km, 588×624 pixels
- Sun position SZA: 30°, SAA: 13°
- Surface albedo data from MODIS
- US standard atmosphere
- ICON clouds (3D liquid and ice water content fields)
- Effective radii parameterized following Bugliaro et al. 2011
- Optical properties:
  - liquid water clouds: Mie
  - ice water clouds: general habit mixture; Yang et al. 2013, Baum et al. 2014

Statistics of **synthetic data** can be compared to real satellite observations to verify whether clouds are realistic.

# Simulated reflectance image for LES scene

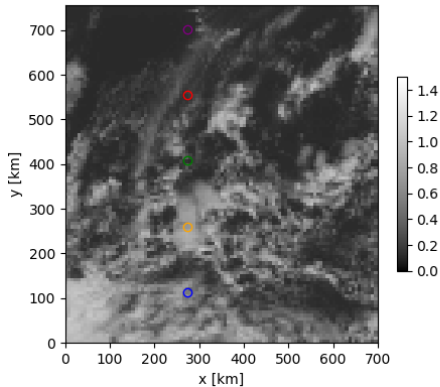


# Sentinel-5 reflectance simulation for LES scene

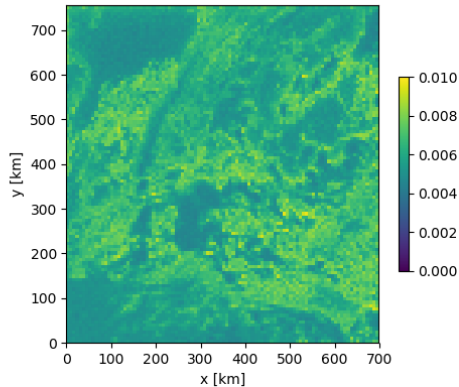
- Spectral range: 400-500 nm (0.2 nm resolution)
- Spatial resolution approx. 7 km,  $98 \times 104$  pixel
- Nadir view
- Sun position SZA:  $30^\circ$ , SAA:  $13^\circ$
- Surface albedo: 0.05
- Molecular optical thickness profiles provided by BIRA
- $\text{NO}_2$ -profile: European polluted
- Layer-AMF calculated at 460 nm
- ICON clouds (3D liquid and ice water content fields, spatial resolution approx. 1.2 km)

# Reflectance simulation

spatial res. approx.  $7 \times 7 \text{ km}^2$

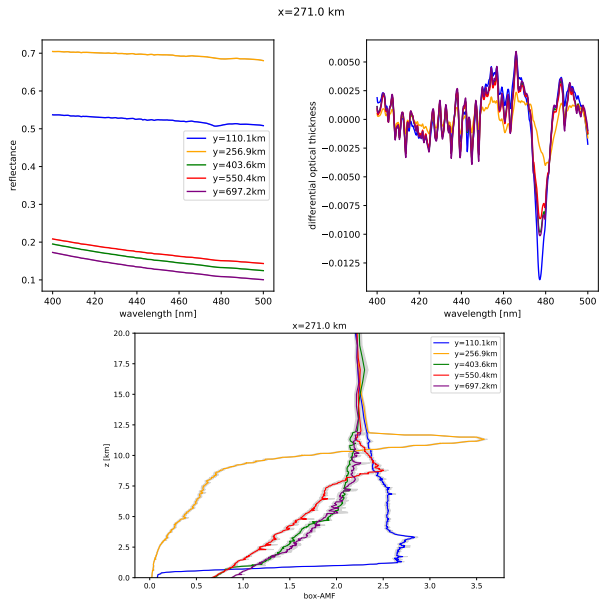


relative standard deviation

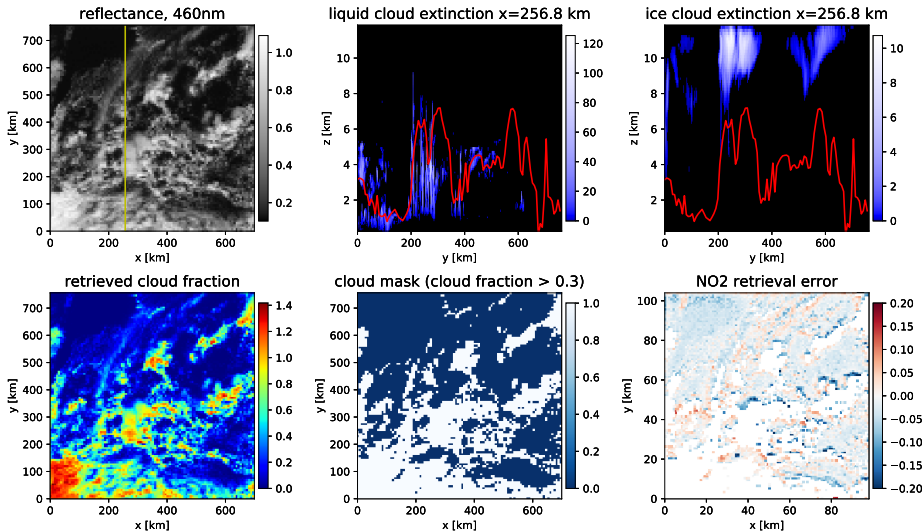




# Reflectance spectra, DOD, and layer-AMF

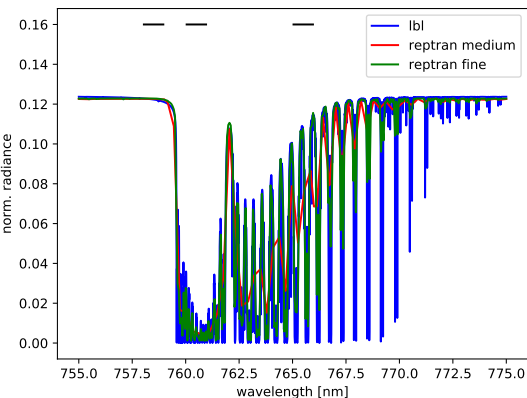


# BIRA-NO2 retrieval ( $O_2-O_2$ )



Largest retrieval errors in cloud shadows

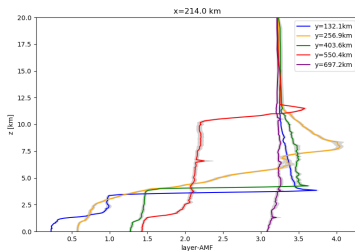
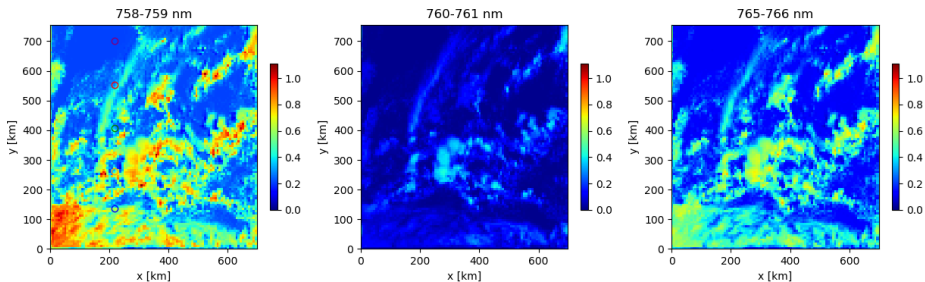
# O<sub>2</sub>A band simulations



- Line-by-line (ARTS, Eriksson et al. 2011)
- REPTRAN absorption parameterization (Gasteiger et al., 2014)
- FRESCO cloud algorithm uses averages over bands 758–759 nm, 760–761 nm, 765–766 nm.

REPTRAN (fine spectral resolution) accuracy sufficient to calculate band averages, saves storage memory and CPU time

# Reflectance and layer-AMF simulations in O<sub>2</sub>A band

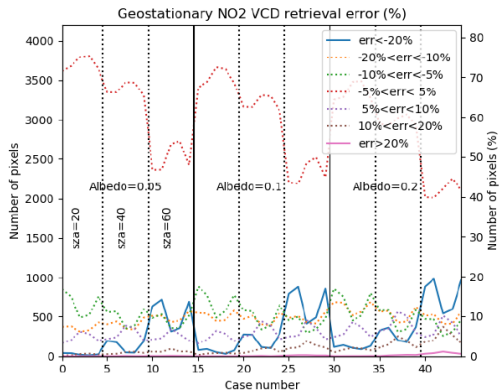


# Synthetic dataset for LES cloud scene

	Geostationary Orbit	Low Earth Orbit
solar zenith angles [°]	20,40,60	20,40,60
solar azimuth angles [°]	-90, 45,0,45,90	13, 353
sensor viewing zenith angle [°]	58.3	0,20,60
sensor viewing azimuth angle [°]	196.3	109.5, 281.7
surface albedo	0,0.05,0.2, (0.5 for O <sub>2</sub> A band)	

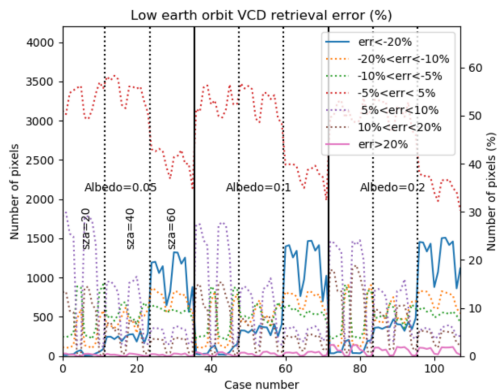
**Table:** Representative sun positions, sensor viewing directions and surface albedos included in synthetic dataset. 45 combinations for GEO and 108 for LEO.

# NO<sub>2</sub> retrieval error statistics - GEO



- average medium bias: -0.9%
- between 61% (high albedo) and 93% (low albedo) of the retrieved NO<sub>2</sub> VCD are within 10% of the “true” column
- number of pixels with differences < -20% increase from about 0.2% for low albedo and high sun to up to 22% for high albedo and low sun
- The variation within each SZA interval is due to different SAA

# NO<sub>2</sub> retrieval error statistics - LEO



- average median bias: -0.5%
- between 53% (high albedo) and 89% (low albedo) of the retrieved NO<sub>2</sub> VCD are within 10% of the actual column
- number of pixels with differences < -20% increase from about 0.1% for low albedo and high sun to up to 26% for high albedo and low sun
- Within each SZA interval the SAA angle takes two values, results are similar for these two angles
- largest differences within each SZA interval are found when VZA is large (60°)

# Summary

- One-dimensional simulations
  - Ensure that MYSTIC and LIDORT agree for clear-sky cases
  - Test cloud correction algorithm ( $O_4$  and FRESCO) for 1D cloud cases
- Box cloud simulations
  - (3D) Box-airmass factor simulations
  - Simulation of reflectance spectra for clear pixels influenced by near clouds
  - Impact of solar zenith angle, albedo, cloud optical thickness, cloud geometrical thickness and cloud bottom height on  $NO_2$  retrieval error
- Comprehensive synthetic dataset with LES cloud input for VIS and  $O_2A$ 
  - Simulated scene covers all typical cloud types for central Europe
  - Sub-pixel cloud inhomogeneity included
  - Quantification of  $NO_2$  retrieval error due to cloud scattering