

# Testing data assimilation methods on convective-scale dynamics

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## Goals

- Create a model hierarchy where the models are computationally cheap but represent some of the key features of convective-scale phenomena.
- Identify and assess promising data assimilation algorithms intended for this scale
- Use the results of toy model experiments to predict the behavior in full models

## Models and experiments developed for testing the suitability:

- Stochastic cloud model
- Modified shallow water equations (presented here)
- Idealized NWP system experiments

## Data assimilation algorithms tested:

- Ensemble Transform Kalman Filter (ETKF)
  - Particle Filter (Sequential Importance Resampling)
  - Efficient particle filter (with nudging)
- All methods with localization and observation averaging

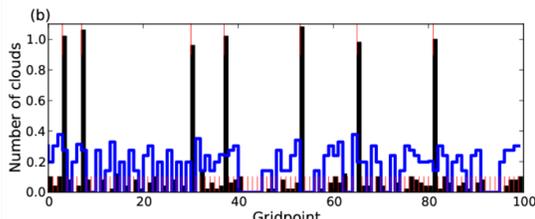
## Stochastic cloud model

### Non-Gaussianity

- Discrete number of clouds in each box
- Low density of clouds (e.g. 0.1)
- No spatial correlation between grid points

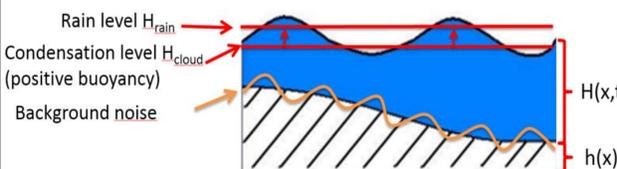
### Nonlinearity

- Clouds appear and disappear randomly
- Poisson birth-death process with instantaneous birth/death
- Probability of death gives average lifetime



Assimilation with ETKF and ensemble size 50. Red vertical lines are position of clouds. ➢ All clouds are assimilated correctly. ➢ Lots of spurious clouds in the ensemble mean.

## Modified shallow water model



1D Shallow water model plus an additional equation for rain. Velocity equation is modified to initiate formation of clouds.

Momentum equation:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + \frac{\partial(\phi + r)}{\partial x} = K \frac{\partial^2 u}{\partial x^2} + F, \phi = \begin{cases} \phi_c + gH, & Z > H_c \\ g(H + h), & \text{otherwise} \end{cases}$$

Continuity equation:

$$\frac{\partial h}{\partial t} + \frac{\partial(uh)}{\partial x} = K \frac{\partial^2 h}{\partial x^2}$$

Rain equation with advection, production and removal of rain

$$\frac{\partial r}{\partial t} + u \frac{\partial r}{\partial x} = K_r \frac{\partial^2 r}{\partial x^2} - \alpha r - \begin{cases} \beta \frac{\partial u}{\partial x}, & Z > H_r \text{ and } \frac{\partial u}{\partial x} < 0 \\ 0, & \text{otherwise,} \end{cases}$$

### Model settings:

Gravity wave speed = 30m/s,  $H_0=90m$   
 $dx=500m, dt=5s, domain=500km, H_c=90.04, H_r=90.4$

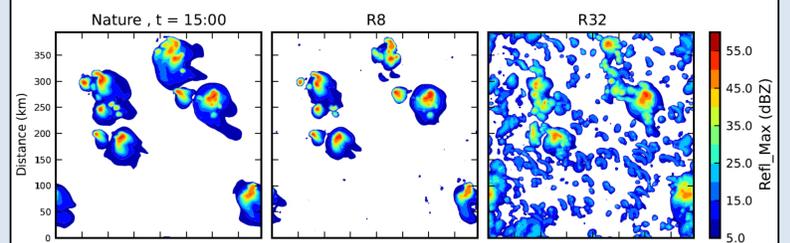
## Idealized NWP System Experiments

### COSMO

Non-hydrostatic, convection-permitting, NWP model  
Domain: 396 x 396 x 20 km, cyclic boundary conditions  
Resolution: 2 km horizontal, 50 vertical levels, 12 s time step  
Initial sounding: 2200 J/kg CAPE, unidirectional shear

### KENDA

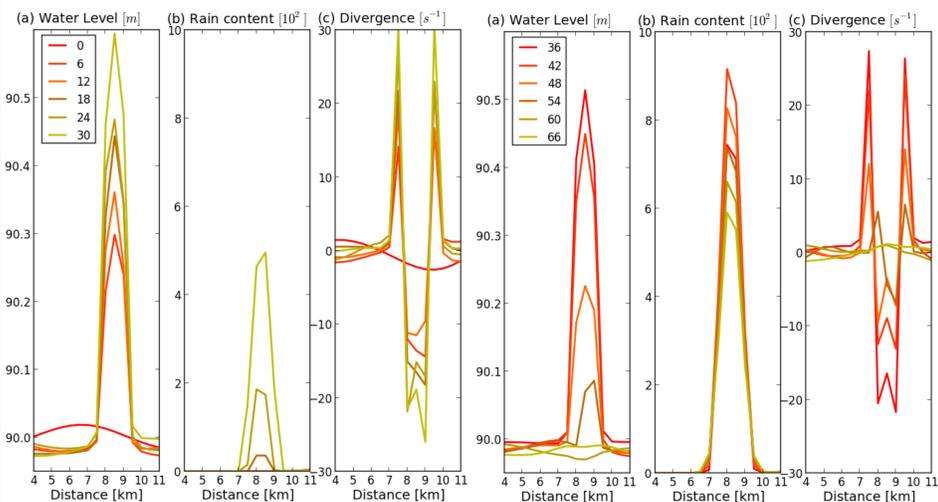
50 members LETKF, initialized with random T and w perturbations. Assimilation every 5 min using Doppler winds, averaged to 4 (or 8) km. Gaspari-Cohn localization radius of 16 (or 32) km.



Analysis after 1 hour of assimilation

- Short-lived clouds are not captured well.
- More spurious clouds in R32 experiment.
- After 3 hours of free forecast R8 and R32 give similar results.

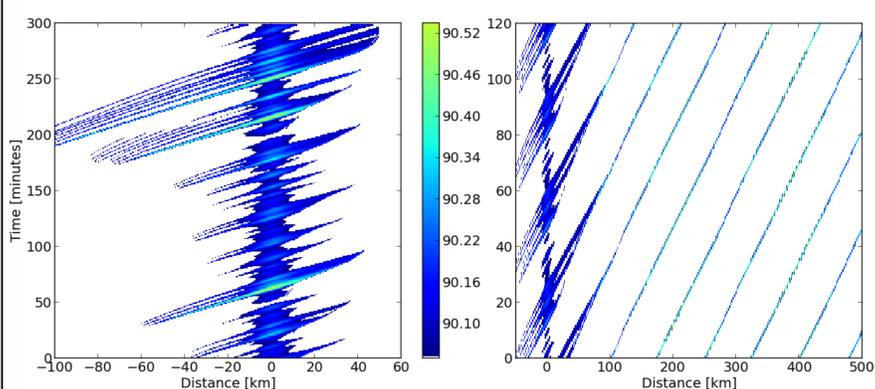
## 1. Life cycle of a convective cloud



Updraft (left) and downdraft (right) phase of a cloud. Different colors correspond to different times in minutes.

- When the water level reaches the cloud threshold, the updraft starts.
- At the water level of 90.4 rain starts to be produced.
- When rain reaches the maximum, the cloud is forced downward again.

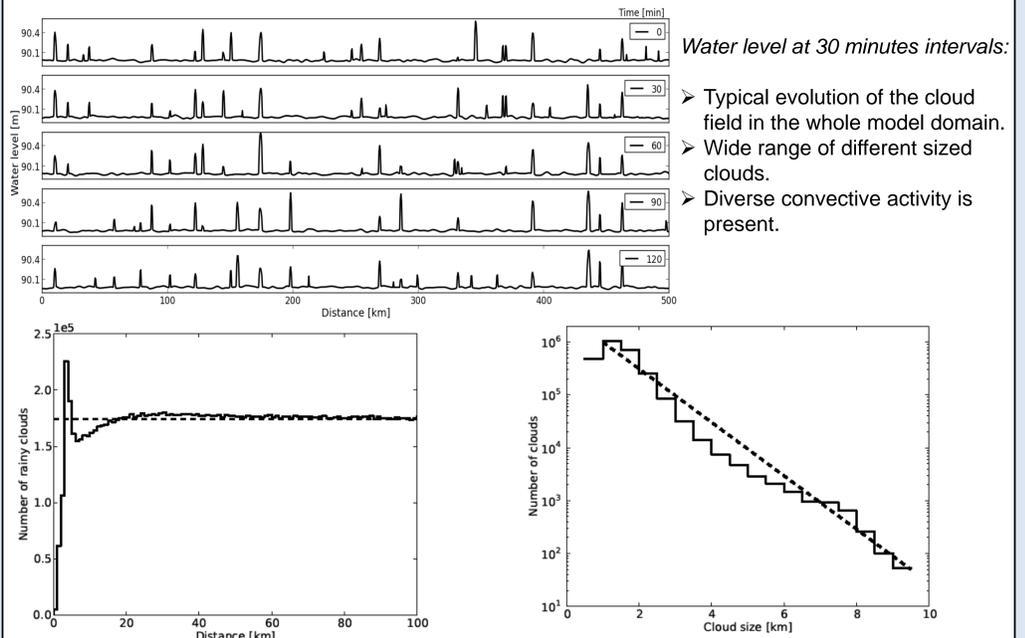
## 3. Orographically triggered convection



Hovmöller diagrams for a mean wind of 30 m/s (left) and 40 m/s (right). A mountain with half-width 10 km and a height of 0.2 m is located at distance 0 km.

- In case of a weak mean wind, the clouds are not able to move past the orography
- In case of a stronger mean wind, a sequence of clouds is built and propagates downstream.
- Comparable to simulations of flow over a ridge (Chu and Lin, 2000)

## 2. Statistics of convection



Water level at 30 minutes intervals:

- Typical evolution of the cloud field in the whole model domain.
- Wide range of different sized clouds.
- Diverse convective activity is present.

Distribution of the distance between different precipitating clouds.

- Maximum at 3.5 km, minimum around 7 km.
- Farther away, the distribution is random.

Logarithmic cloud size distribution for clouds with a height above 90.04.

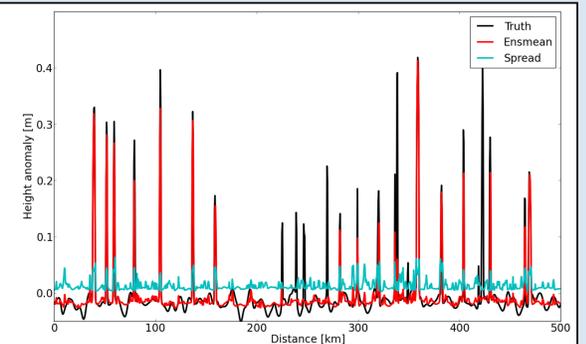
- Average cloud size is 3.4 km.
- Mean number of clouds in domain is 14.9.
- Therefore 5% of the grid points are cloudy.

## 4. Data assimilation

- Radar observations and radial wind as synthetic observations.
- More thorough analysis of different observation configurations ongoing.

### Preliminary Results:

- A good fraction of clouds is assimilated when observing all grid points and using quite a small observation error.
- Fastest evolving clouds not captured.
- Ensemble is underdispersive.



Height analysis after 10 LETKF cycles with observations every 5 minutes and 20 ensembles. All fields are observed at every grid point.

## References

- Chu, C.-M. and Y.-L. Lin, 2000: Effects of Orography on the Generation and Propagation of Mesoscale Convective Systems in a Two-Dimensional Conditionally Unstable Flow. – J. Atmos. Sci. **57**, 3817-3837
- Craig, G. C. and Würsch, M., 2012: The impact of localization and observation averaging for convective-scale data assimilation in a simple stochastic model. Q.J.R. Meteorol. Soc., doi: 10.1002/qj.1980
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